

Unstructured-mesh based non-oscillatory-forward-in-time integration for atmospheric flows

J. Szmelter

Loughborough University, UK

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Generalised Nonhydrostatic Model

Smolarkiewicz et. al. JCP 2016

$$\frac{\partial \mathcal{G}_\rho}{\partial t} + \nabla \cdot (\mathcal{G}_\rho \mathbf{v}) = 0$$

$$\frac{\partial \mathcal{G}_\rho \theta'}{\partial t} + \nabla \cdot (\mathcal{G}_\rho \mathbf{v} \theta') = -\mathcal{G}_\rho \left(\tilde{G}^T \mathbf{u} \cdot \nabla \theta_a - \mathcal{H} \right)$$

$$\frac{\partial \mathcal{G}_\rho \mathbf{u}}{\partial t} + \nabla \cdot (\mathcal{G}_\rho \mathbf{v} \otimes \mathbf{u}) = -\mathcal{G}_\rho \left(\Theta \tilde{G} \nabla \varphi + g \Upsilon_B \frac{\theta'}{\theta_b} + \mathbf{f} \times (\mathbf{u} - \Upsilon_C \mathbf{u}_a) - \mathcal{M}'(\mathbf{u}, \mathbf{u}, \Upsilon_C) - \mathcal{D} \right)$$

$$\rho := \left[\rho(\mathbf{x}, t), \frac{\rho_b(z)\theta_b(z)}{\theta(\mathbf{x}, t)}, \rho_b(z) \right] \quad \varphi := [c_p \theta_0 \pi', c_p \theta_0 \pi, c_p \theta_b \pi] \quad \Theta := \left[\frac{\theta}{\theta_0}, \frac{\theta}{\theta_0}, 1 \right] \quad \Upsilon_B := \left[\frac{\theta_b(z)}{\theta_a(\mathbf{x})}, \frac{\theta_b(z)}{\theta_a(\mathbf{x})}, 1 \right] \quad \Upsilon_C := \left[\frac{\theta}{\theta_a(\mathbf{x})}, \frac{\theta}{\theta_a(\mathbf{x})}, 1 \right]$$

[compressible, pseudo-incompressible, anelastic]

$$\varphi = c_p \theta_0 \left[\left(\frac{R_d}{p_0} \rho \theta \right)^{R_d/c_v} - \pi_a \right]. \quad \text{Ideal gas law for compressible option}$$

NFT based Nonhydrostatic Models

Non-oscillatory Forward-in-Time (NFT)

$$\frac{\partial G\Phi}{\partial t} + \nabla \cdot (\mathbf{V}\Phi) = G\mathcal{R}$$

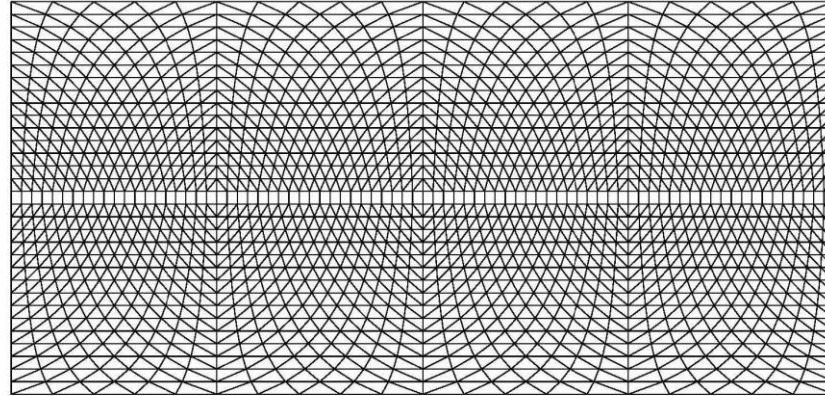
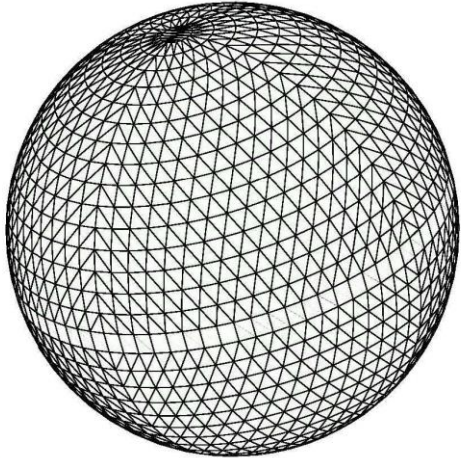
$$\Phi_i^{n+1} = \mathcal{A}_i \left(\Phi^n + \frac{1}{2}\delta t\mathcal{R}, \mathbf{V}^{n+\frac{1}{2}}, G^n, G^{n+1} \right) + \frac{1}{2}\delta t\mathcal{R}^{n+1}$$

- Semi-implicit
- Multidimensional Positive Definite Advection Transport Algorithm (MPDATA)
- Krylov solver

For details see:

Kühnlein et. al. GMD 2019 for global model (compressible) FVM (Finite Volume Module) of IFS
Szmelter et. al. JCP 2019 for local model (anelastic)

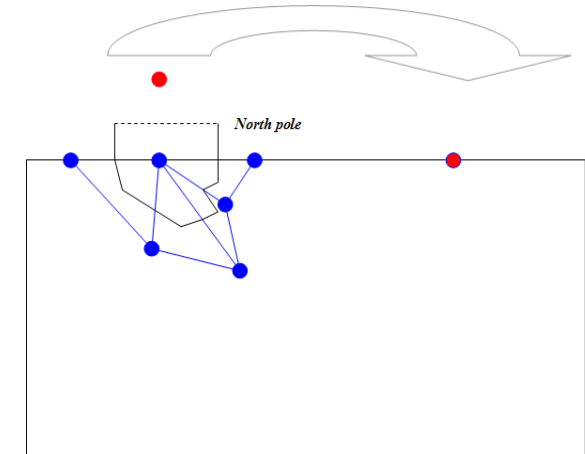
Global flows - FVM: discrete operators



Periodic boundaries



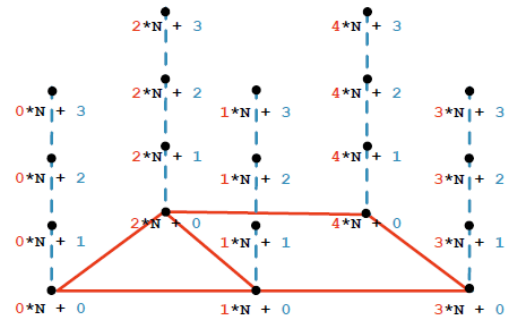
Pole treatment



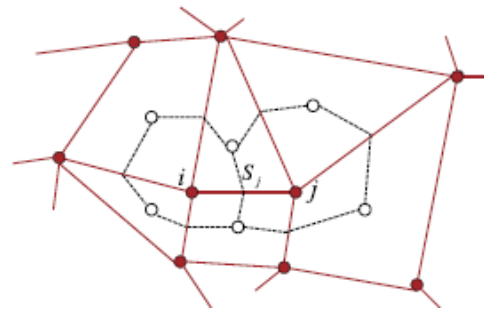
Szmelter & Smolarkiewicz JCP 2010

Octahedral mesh

IFS grid



Uses a co-located data arrangement for all prognostic variables



$$\int_{\Omega} \nabla \cdot \mathbf{A} = \int_{\partial\Omega} \mathbf{A} \cdot \mathbf{n} = \frac{1}{\mathcal{V}_i} \sum_{j=1}^{l(i)} A_j^{\perp} S_j$$

dual volume: \mathcal{V}_i , face area: S_j

Generalised Conjugate Residual (k) scheme

For any initial guess, ϕ^0 , set $r^0 = \mathcal{L}(\phi^0) - R$, $p^0 = \mathcal{P}^{-1}(r^0)$; then iterate:

For $n = 1, 2, \dots$ until convergence

Smolarkiewicz & Margolin 2000

$$\sum_{I=1}^M \frac{\partial}{\partial x^I} \left(\sum_{J=1}^M C^{IJ} \frac{\partial \phi}{\partial x^J} + D^I \phi \right) - A\phi = R$$

$$\mathcal{L}(\phi) - R = 0 .$$

$$\frac{\partial^k \mathcal{P}(\phi)}{\partial \tau^k} + \frac{1}{T_{k-1}(\tau)} \frac{\partial^{k-1} \mathcal{P}(\phi)}{\partial \tau^{k-1}} + \dots + \frac{1}{T_1(\tau)} \frac{\partial \mathcal{P}(\phi)}{\partial \tau} = \mathcal{L}(\phi) - R$$

for $\nu = 0, \dots, k - 1$

$$\beta = -\frac{\langle r^\nu \mathcal{L}(p^\nu) \rangle}{\langle \mathcal{L}(p^\nu) \mathcal{L}(p^\nu) \rangle},$$

$$\phi^{\nu+1} = \phi^\nu + \beta p^\nu,$$

$$r^{\nu+1} = r^\nu + \beta \mathcal{L}(p^\nu),$$

exit if $\|r^{\nu+1}\| \leq \epsilon$,

$$e = \mathcal{P}^{-1}(r^{\nu+1}),$$

evaluate $\mathcal{L}(e)$

for $l = 0, \dots, \nu$

$$\alpha_l = \frac{\langle \mathcal{L}(e) \mathcal{L}(p^l) \rangle}{\langle \mathcal{L}(p^l) \mathcal{L}(p^l) \rangle},$$

$$p^{\nu+1} = e + \sum_l^{\nu} \alpha_l p^l,$$

$$\mathcal{L}(p^{\nu+1}) = \mathcal{L}(e) + \sum_l^{\nu} \alpha_l \mathcal{L}(e^l),$$

reset $[\phi, r, e, \mathcal{L}(e)]^k$ to $[\phi, r, e, \mathcal{L}(e)]^0$

Richardson iteration

Take an operator $\mathcal{P} \approx \mathcal{L}$

$$\sum_{I=1}^M \frac{\partial}{\partial x^I} \left(\sum_{J=1}^M C^{IJ} \frac{\partial \phi}{\partial x^J} + D^I \phi \right) - A\phi = R$$

by neglecting cross derivative terms i.e. $I \neq J$

Solve the resulting residual equation:

$$\frac{\partial e}{\partial \tilde{\tau}} = \mathcal{P}(e) - r \quad \Rightarrow \quad e^{\mu+1} = e^{\mu} + \Delta \tilde{\tau} [\mathcal{P}(e^{\mu}) - r^{\nu+1}]$$

Due to the thin shell atmosphere: treat the unstructured horizontal direction explicitly and the structured vertical direction implicitly, i.e.

$$\text{let } \mathcal{P}(e) \equiv \mathcal{P}_H(e^{\mu}) + \mathcal{P}_z(e^{\mu+1})$$

$$(I - \Delta \tilde{\tau} \mathcal{P}_z) e^{\mu+1} = e^{\mu} + \Delta \tilde{\tau} (\mathcal{P}_H(e^{\mu}) - r^{\nu+1}) = \mathcal{R}^{\mu}$$

$$e^{\mu+1} = [I - \Delta \tilde{\tau} \mathcal{P}_z]^{-1} \mathcal{R}^{\mu} \quad \text{Use Thomas Algorithm}$$

Unstructured mesh Jacobi iteration

Solve the residual equation by lagging the off-diagonal \mathcal{P} components:

$$\mathcal{P}(e) = r$$

$$\mathcal{P}(e) + \mathcal{D}(e) - \mathcal{D}(e) = r$$

$$e^{\mu+1} = e^{\mu} - \frac{\mathcal{P}(e^{\mu}) - r^{\nu+1}}{\mathcal{D}}$$

Apply the thin shell atmosphere treatment again:

$$\text{let } \mathcal{P}(e) \equiv \mathcal{P}_H(e^{\mu}) + \mathcal{P}_z(e^{\mu+1})$$

$$\text{For the Helmholtz } 0 = - \sum_{\ell=1}^3 \left(\frac{A_{\ell}^*}{\zeta_{\ell}} \nabla \cdot \zeta_{\ell} \tilde{\mathbf{G}}^T (\tilde{\mathbf{u}} - \mathbf{C} \nabla \varphi') \right) - B^* (\varphi' - \hat{\varphi}')$$

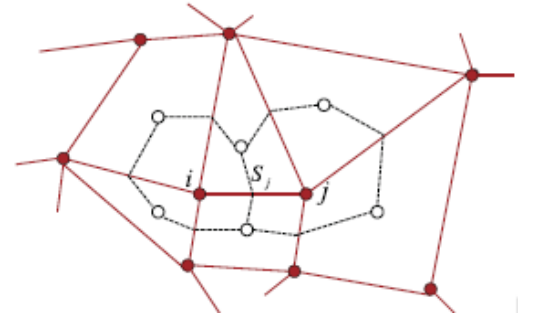
$$\mathcal{D}_{k,i} = - \frac{1}{4V_i} \sum_l \frac{A_{l k,i}^*}{\zeta_{l k,i}} \sum_{j=1}^{nbrs} \frac{\zeta_{l k,j}}{V_j} \left(\mathcal{S}_{x_j}^2 (\tilde{\mathbf{G}}^T \mathbf{C})_{xx k,j} + \mathcal{S}_{y_j}^2 (\tilde{\mathbf{G}}^T \mathbf{C})_{yy k,j} \right)$$

$$\left(I + \frac{\mathcal{P}_z}{\mathcal{D}} \right) \hat{e}^{\mu+1} = e^{\mu} - \frac{\mathcal{P}_H(e^{\mu}) - r^{\nu+1}}{\mathcal{D}} \equiv \frac{\mathcal{R}^{\mu}}{\mathcal{D}}$$

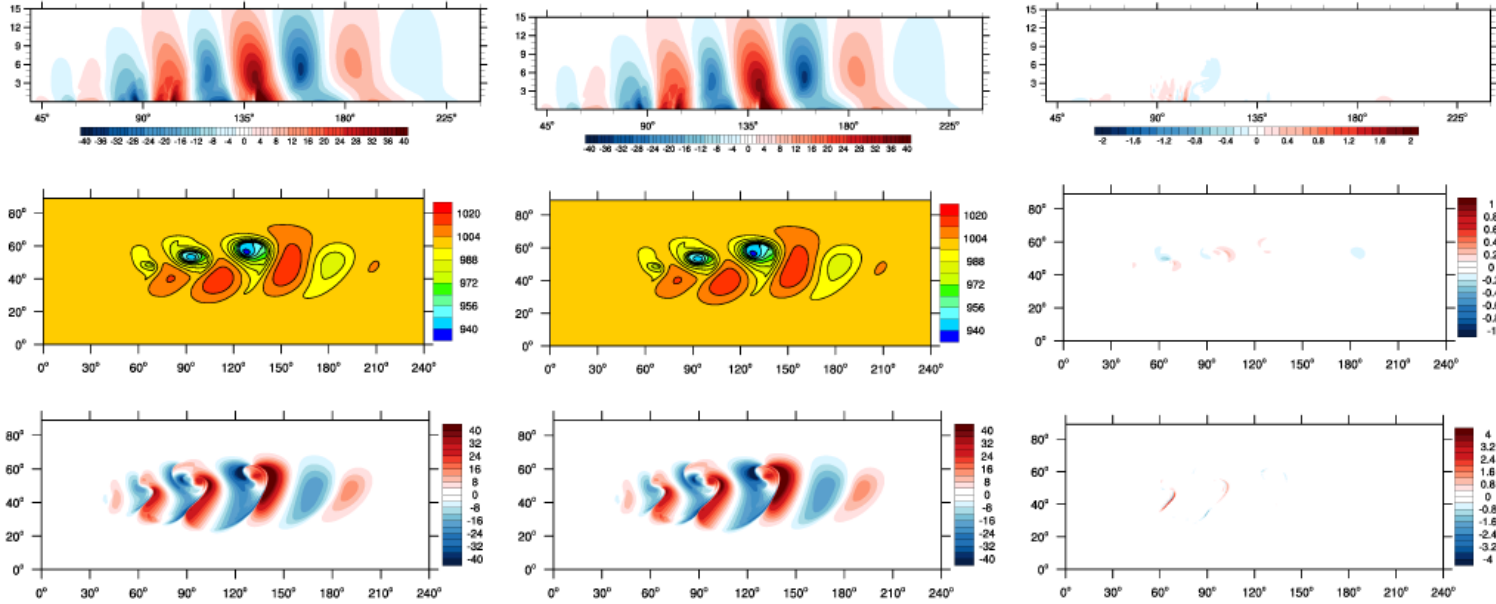
$$\hat{e}^{\mu+1} = [\mathcal{D} + \mathcal{P}_z]^{-1} \mathcal{R}^{\mu} \quad \text{Use Thomas Algorithm}$$

Jacobi requires some solution relaxation:

$$e^{\mu+1} = (1 - \omega) e^{\mu} + \omega \hat{e}^{\mu+1}, \quad \omega \approx \frac{2}{3}$$



Dry baroclinic instability, O360 mesh, Jacobi (left), Richardson (centre), Difference (right)



Vertical velocity at 53° N latitude

Upper Hemisphere
surface pressure

Upper Hemisphere
surface meridional velocity

Table 1: 10 days, dry baroclinic instability test with 60 vertical levels, computed on the Cray XC30 at ECMWF (Dynamical core only).

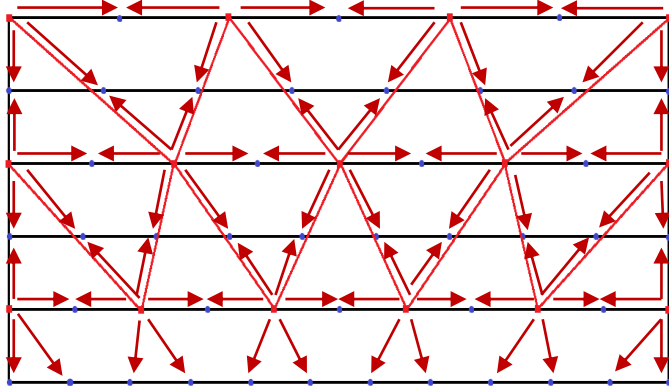
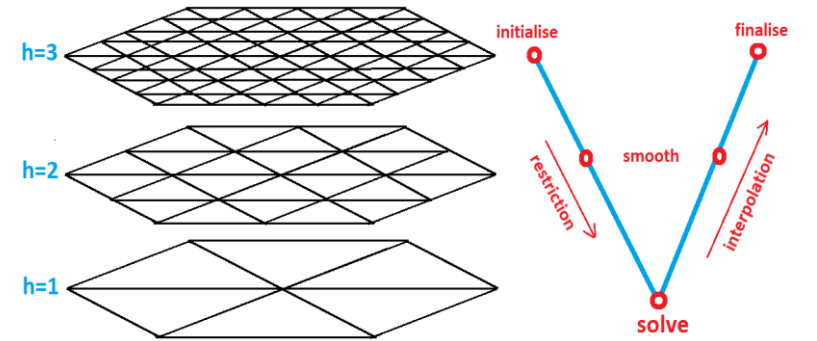
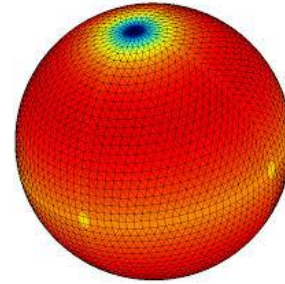
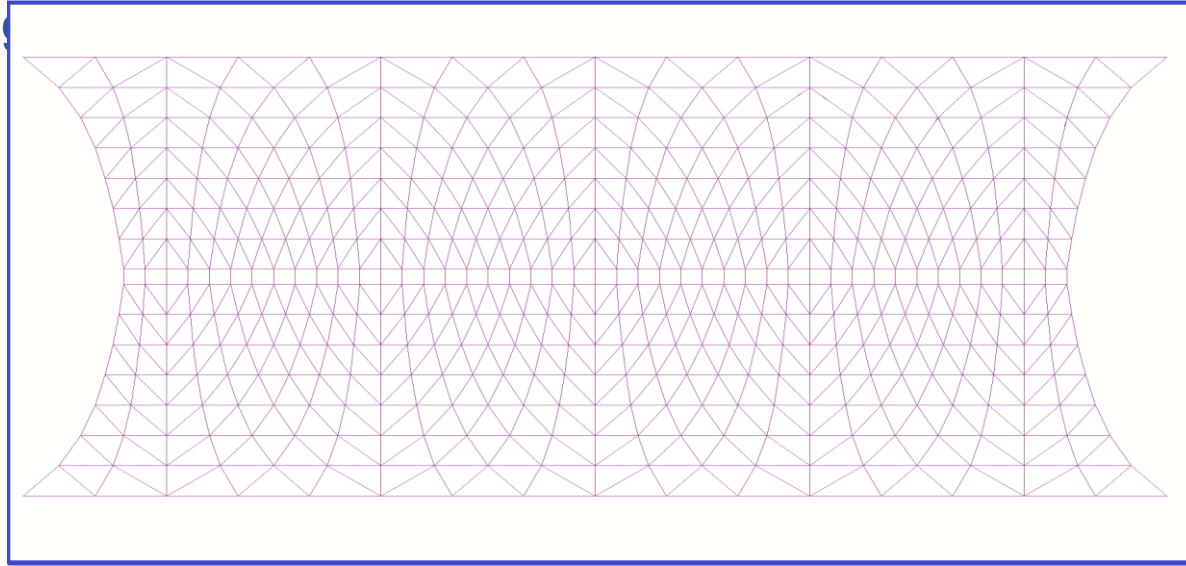
Grid	Richardson runtime	Richardson convergence	Jacobi runtime	Jacobi convergence	Tasks x Threads	Total cores
<i>O180</i>	2897 s	14	1540 s	1	36 x 1	36
<i>O360</i>	3700 s	22	1374 s	1	108 x 6	648
<i>O720</i>	11296 s	33	3194 s	2	360 x 6	2160
<i>O1280</i>	35884 s	47	7976 s	2	720 x 6	4320
<i>O1800</i>	46670 s	56	8675 s	2	2400 x 6	14400

Jacobi provides ca 2x
acceleration on coarse meshes
expanding to ca 5-6 x
acceleration on finer meshes

Preconditioners use 3 iterations

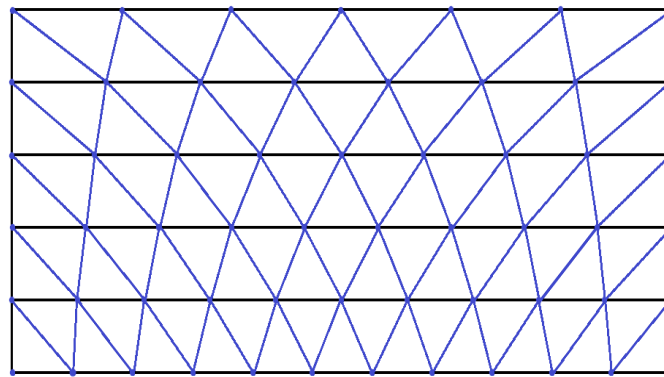
Unstructured mesh multigrid preconditioner in the horizontal

Parallel restriction, prolongation and Atlas mesh



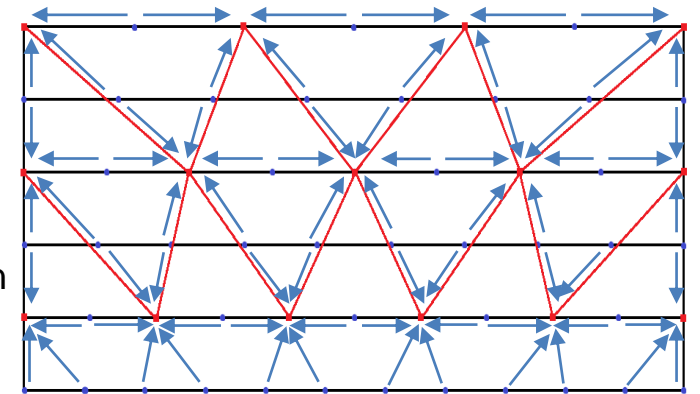
Coarse mesh

interpolation



Fine mesh

restriction

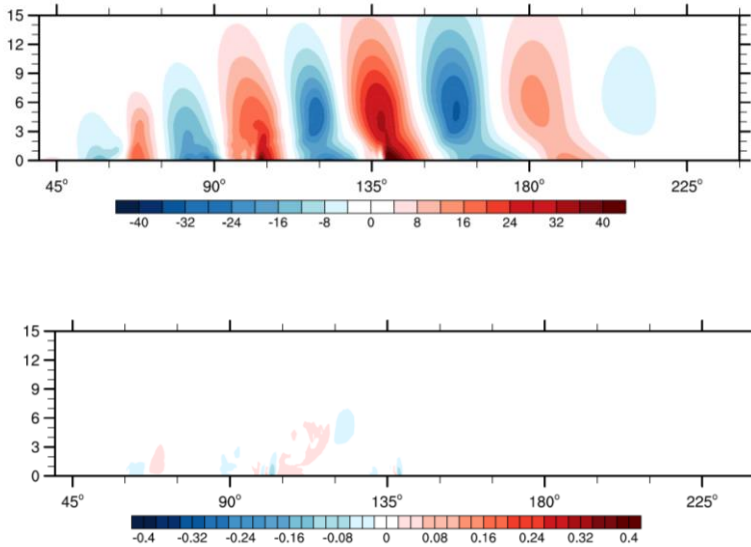


Coarse mesh

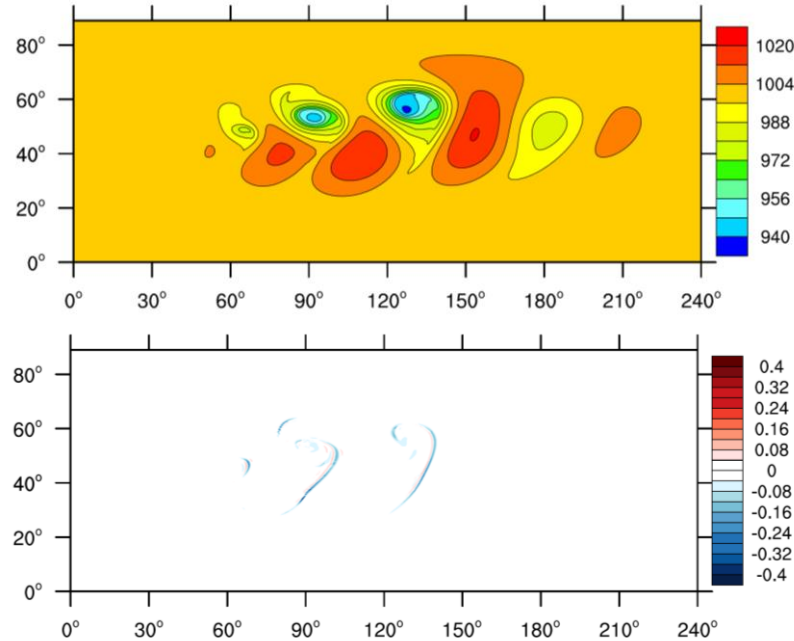
Baroclinic Instability Testcase

O360 plots for day 10 of dry baroclinic instability test: Multigrid (top), Difference vs. Richardson (bottom).

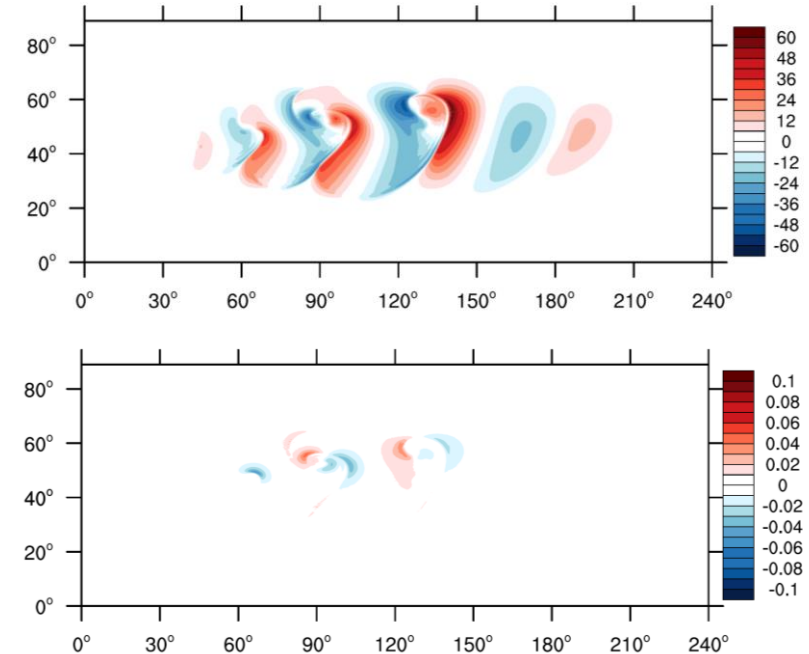
vertical velocity at the 53° N;



Upper Hemisphere surface Pressure



Upper Hemisphere surface meridional velocity



Grid	Richardson	Multigrid	Speedup	Task x thread
O360	814 s	403 s	2.02x	108 x 6 (648 cores)
O720	3001 s	1017 s	2.95x	360 x 6 (2160 cores)
O1280	13586 s	3038 s	4.41x	720 x 6 (4320 cores)
O1800	19544 s	4060 s	4.81x	2400 x 6 (14400 cores)

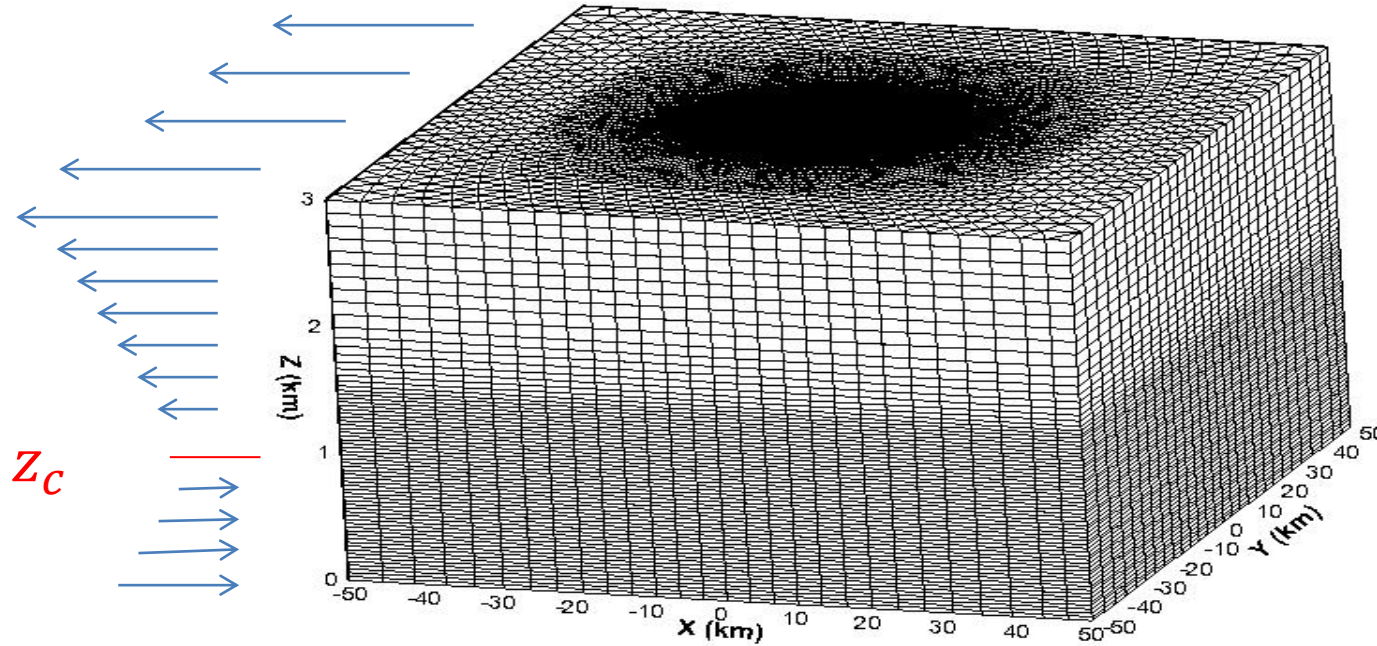
3 iterations

1 Jacobi as a smoother and 3 Jacobi on the coarse grid

10 days, dry baroclinic instability test ~27km Horizontal resolution, with 31 vertically stretched levels at a depth of 45km computed on the Cray XC30 at ECMWF (Dynamical core only)

Local models: The effect of critical levels on stratified flows past an axisymmetric mountain

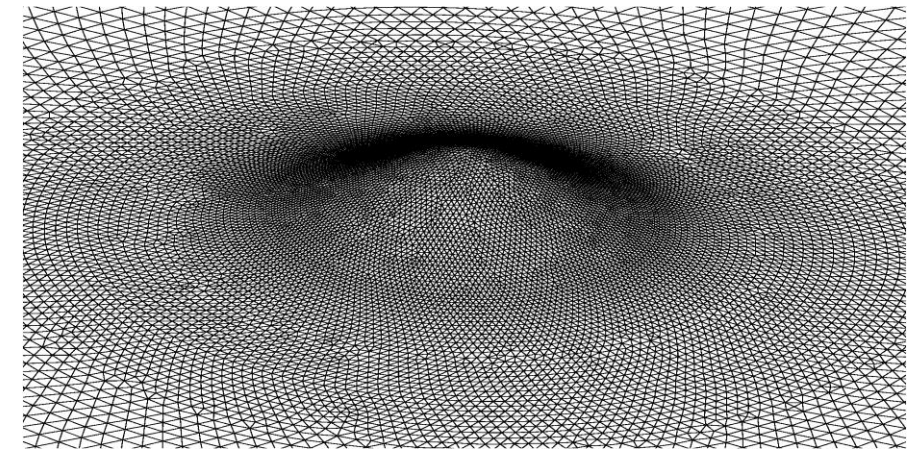
Szmelter et. al. JCP 2015



Prismatic mesh

$$\tilde{z}_{i,k} = \tilde{z}_{i,k-1} + \delta \tilde{z}_k$$

$$\tilde{z}_{i,k} = \tilde{z}_{i,k} \left(1 - \frac{h_i}{H} \right) + h_i$$



Triangular surface mesh

$$U_0 = 10 \text{ m s}^{-1}, \quad U(z) = U_0 \left(1 - \frac{z}{z_c} \right)$$

$$\frac{U_0}{N a} = 0.2, \quad R_i = \left(\frac{N z_c}{U_0} \right)^2 = 1, \quad N = 0.01 \text{ s}^{-1}$$

$$a = 5000 \text{ m}, \quad h(r) = h_0 \left(1 + \frac{r^2}{a^2} \right)^{-\frac{3}{2}}, \quad r \equiv \sqrt{x^2 + y^2}$$

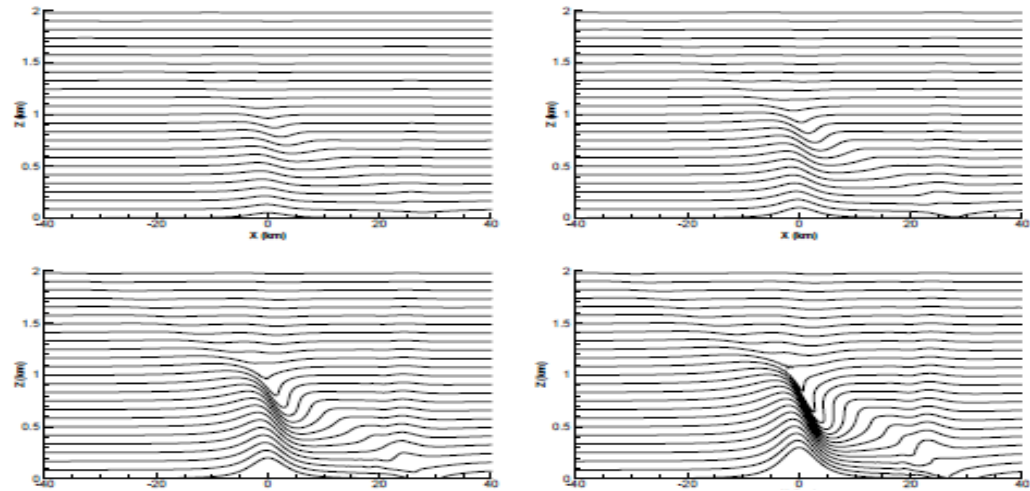


Fig. 12. Isentropes at $T=6$ in $y = 0$ vertical plane for experiments LS2 (left, top) LS3 (right, top), LS4 (left, bottom) and LS5 (right, bottom).

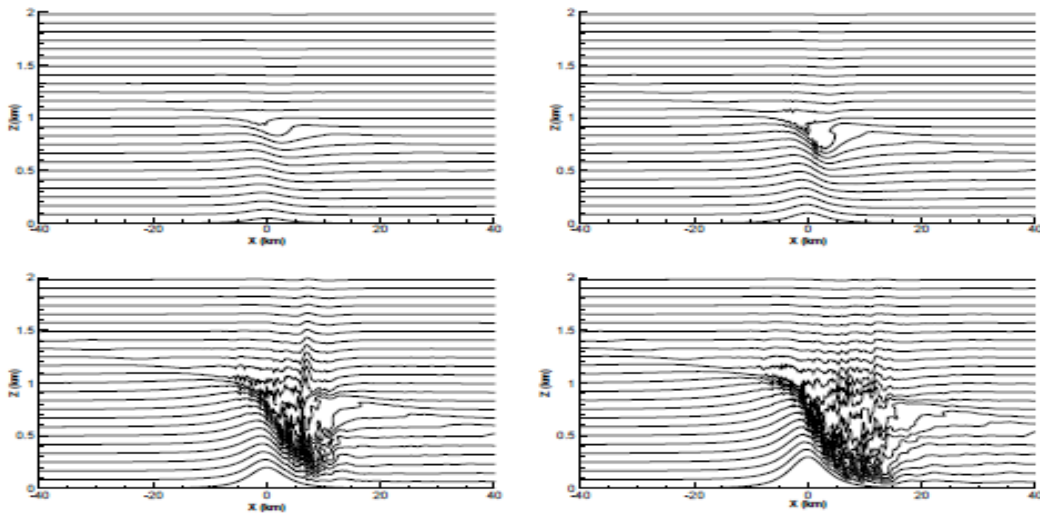
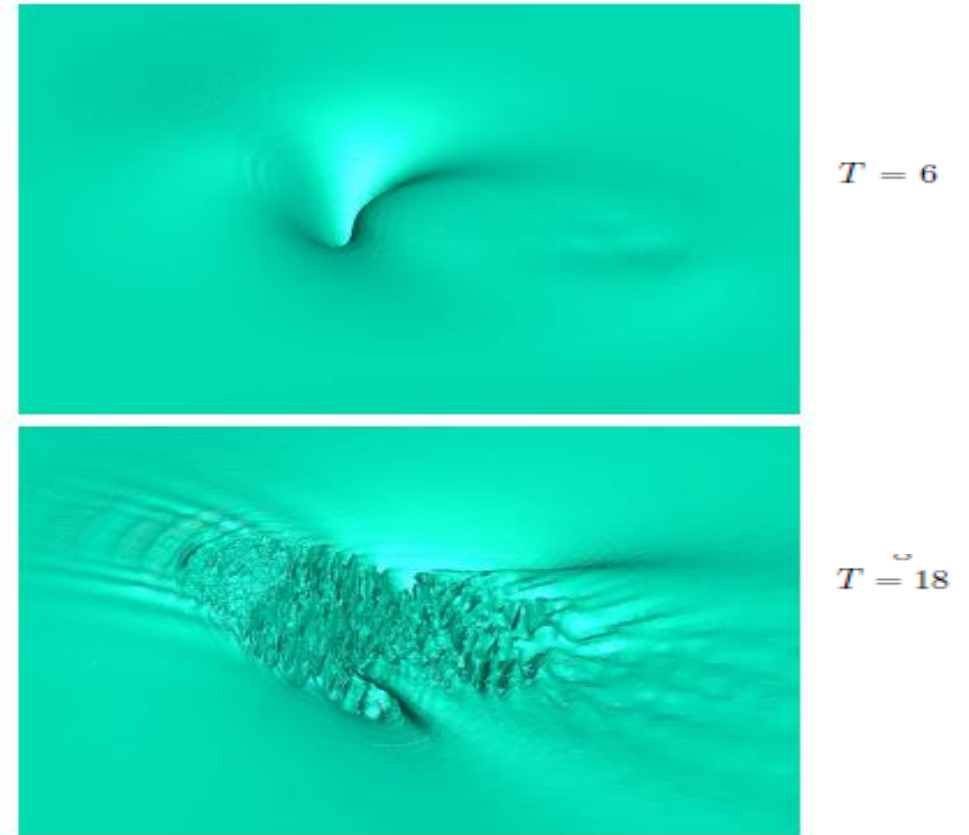


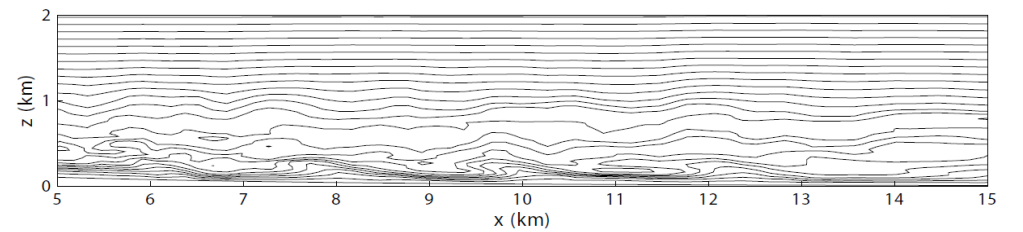
Fig. 13. As in Fig. 12 but at $T=18$.



$T = 6$

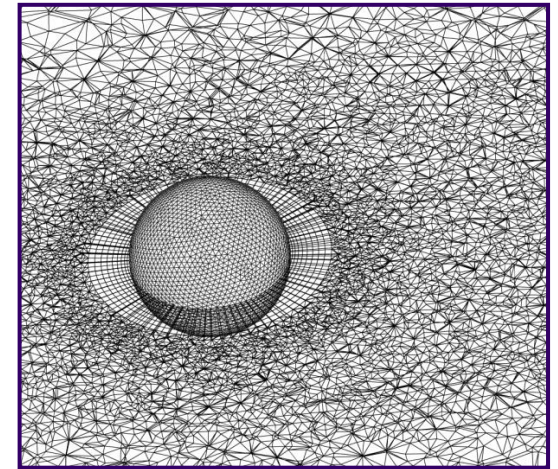
$T = 18$

the isentropes with undisturbed height $z = 0.94z_c$ LS5.



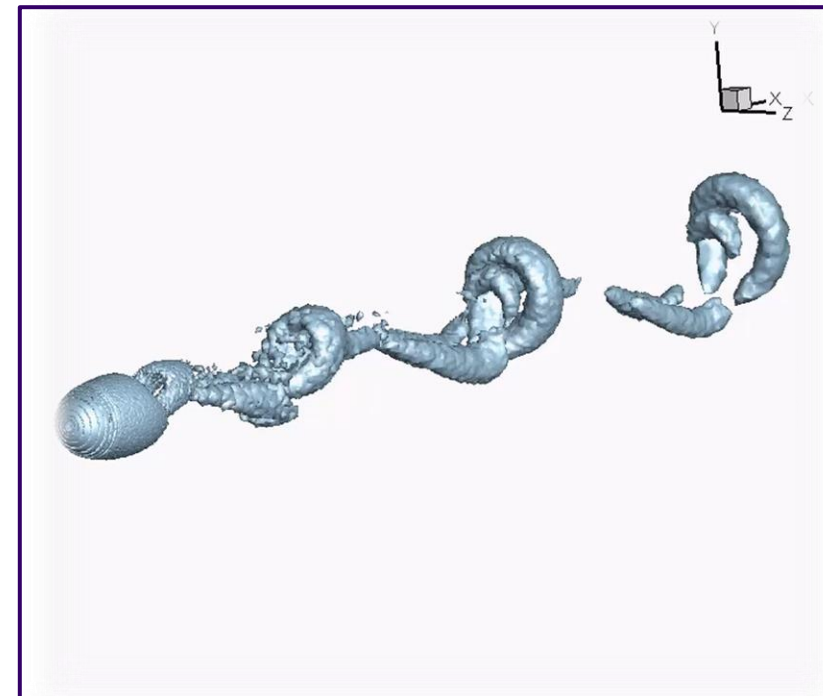
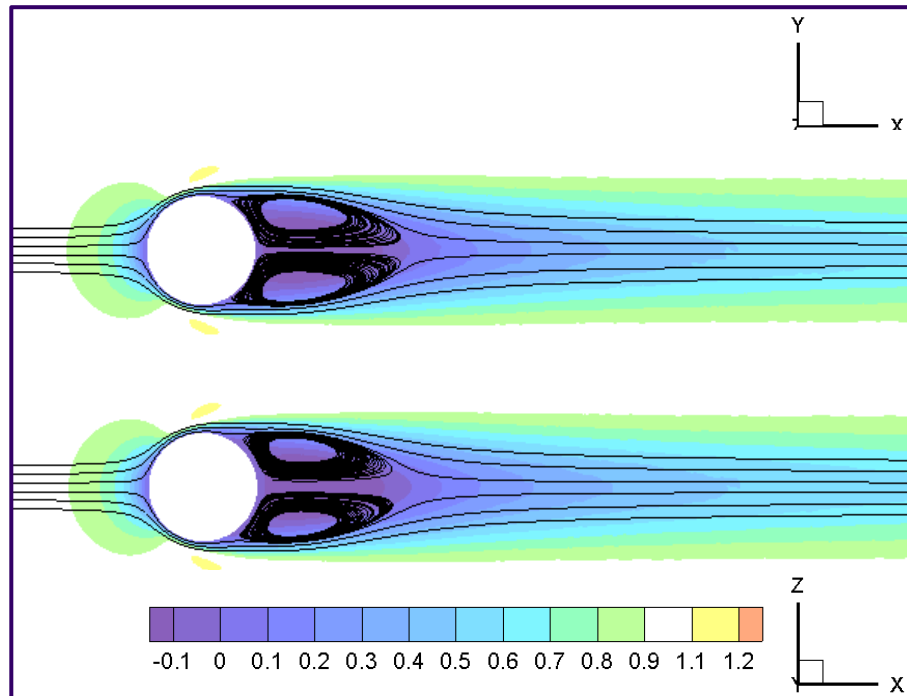
Stratified flows past a sphere

Cocetta et al. PoF 2021



$Re = 200$

$Re = 300$

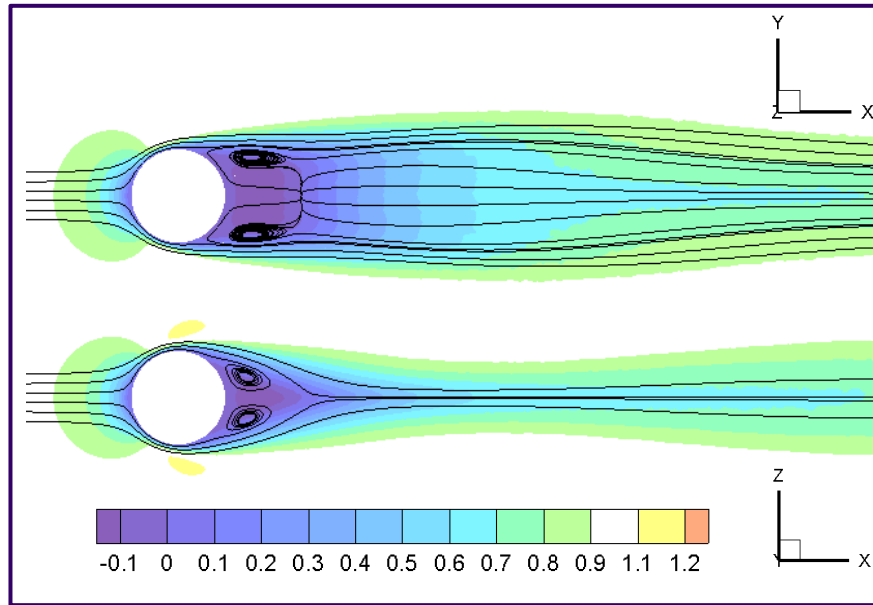


*Q-method
(second invariant
of the deformation
tensor)*

Neutrally-stratified flow $Fr = \infty$

Non-axisymmetric attached vortex regime

$Fr = 2.5$



Re=200

Re=300

Top view



$Fr = 1.66$

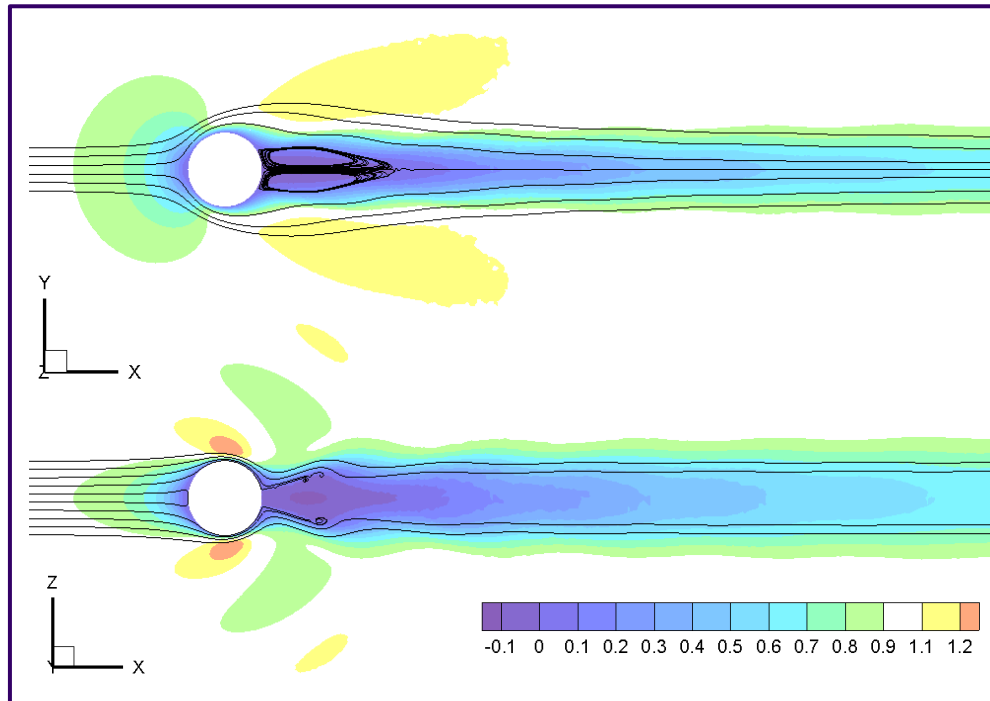
Side view



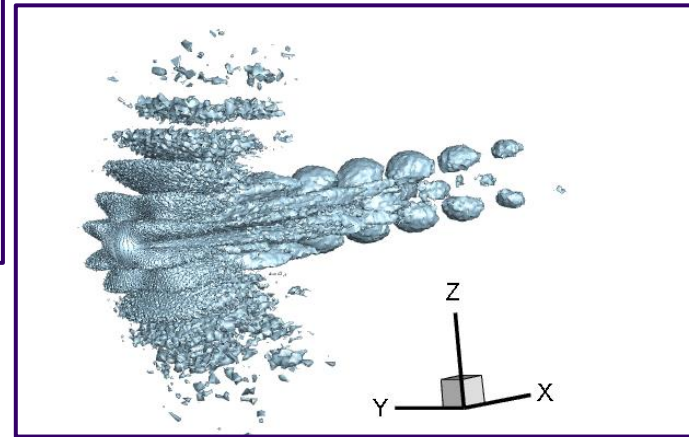
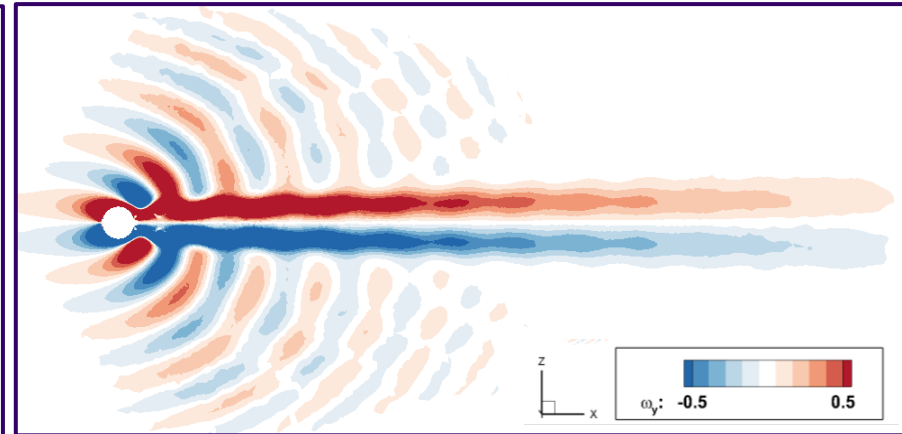
Lee-waves instability regime

$$0.8 \gtrsim Fr \gtrsim 0.6$$

$$Re = 200$$

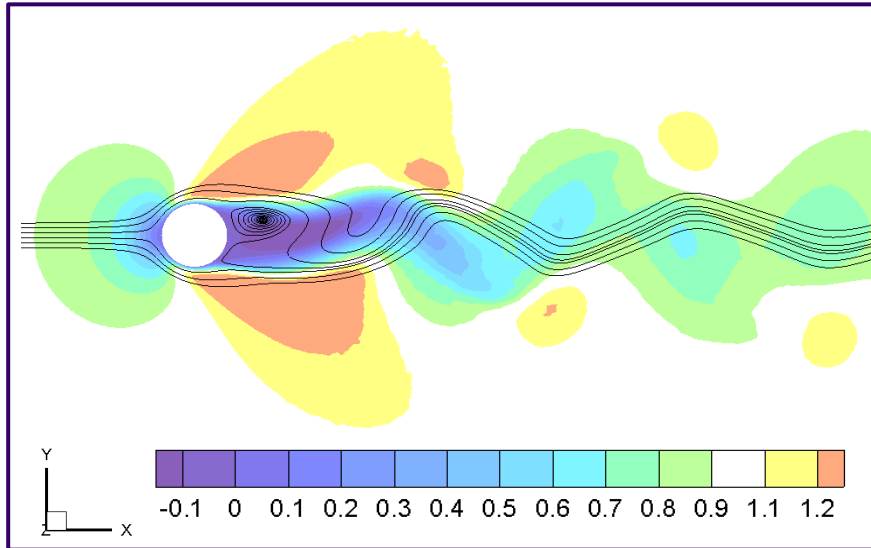


$$Fr = 0.625$$



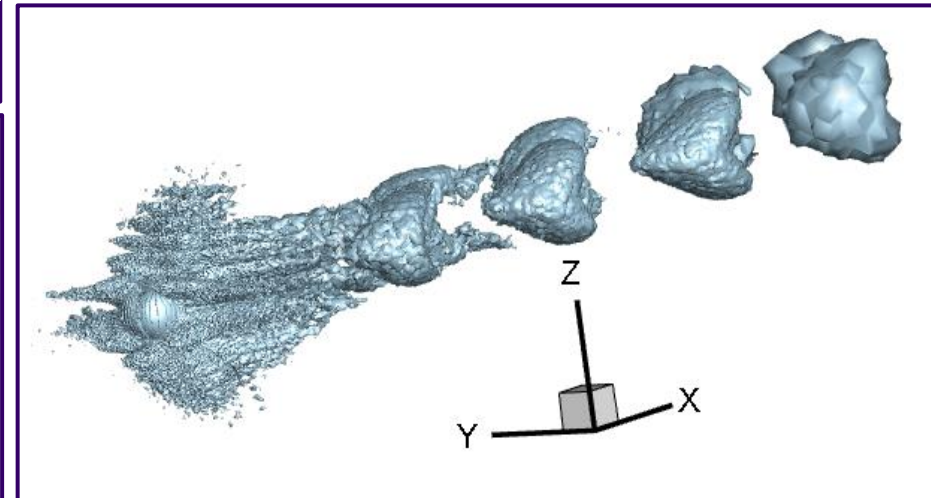
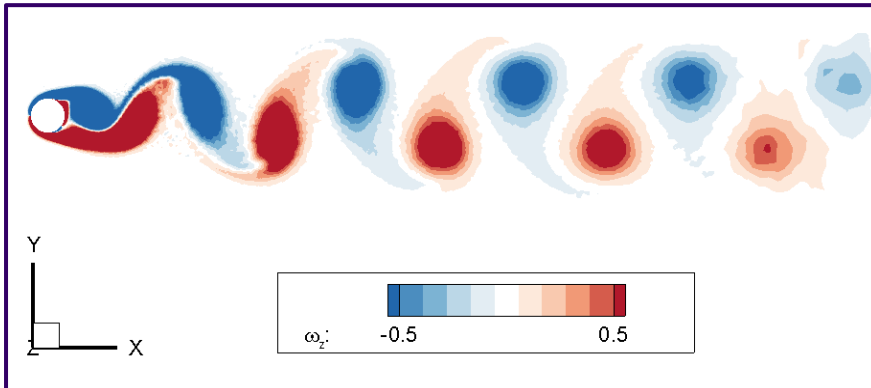
Two-dimensional vortex shedding regime

$$Fr \lesssim 0.6$$



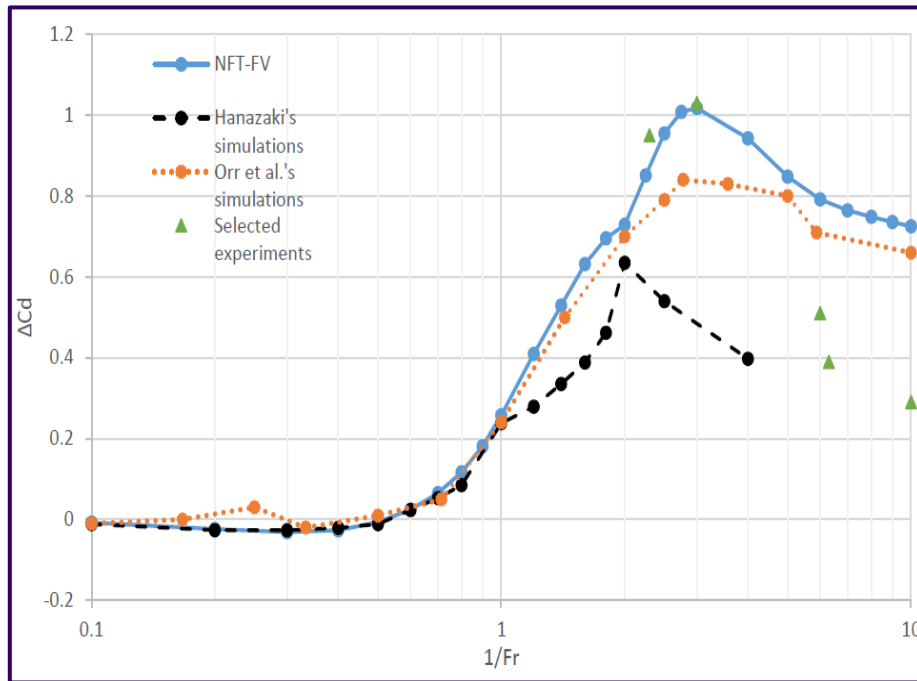
$$Re = 200$$

$$Fr = 0.44$$

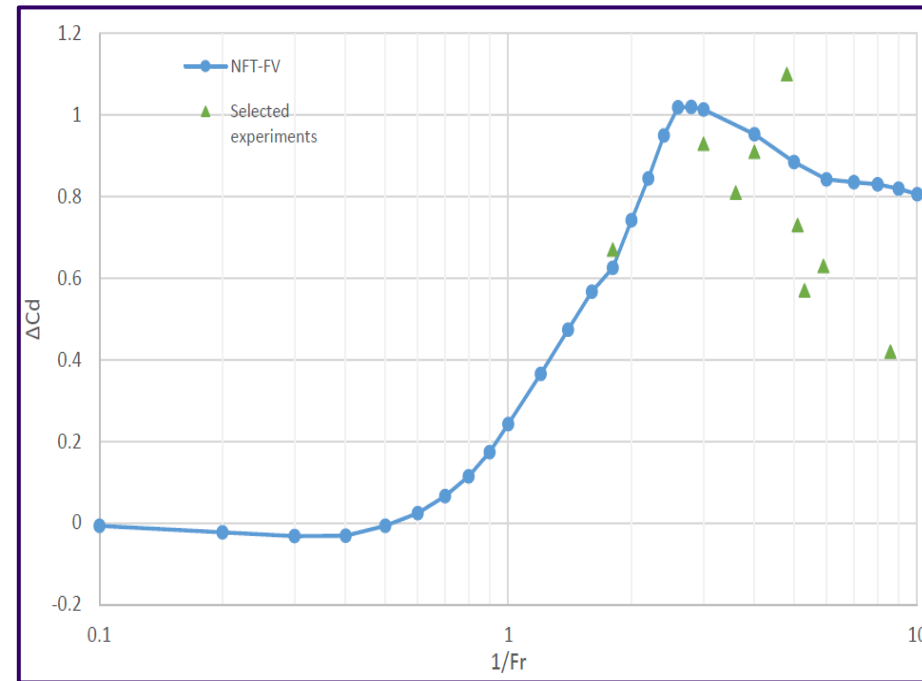


Study of drag coefficient

$Re = 200$



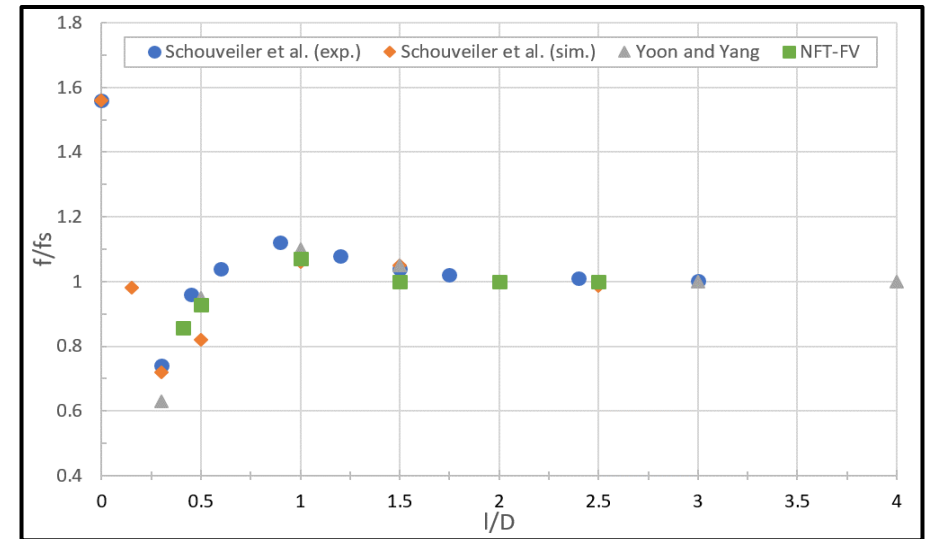
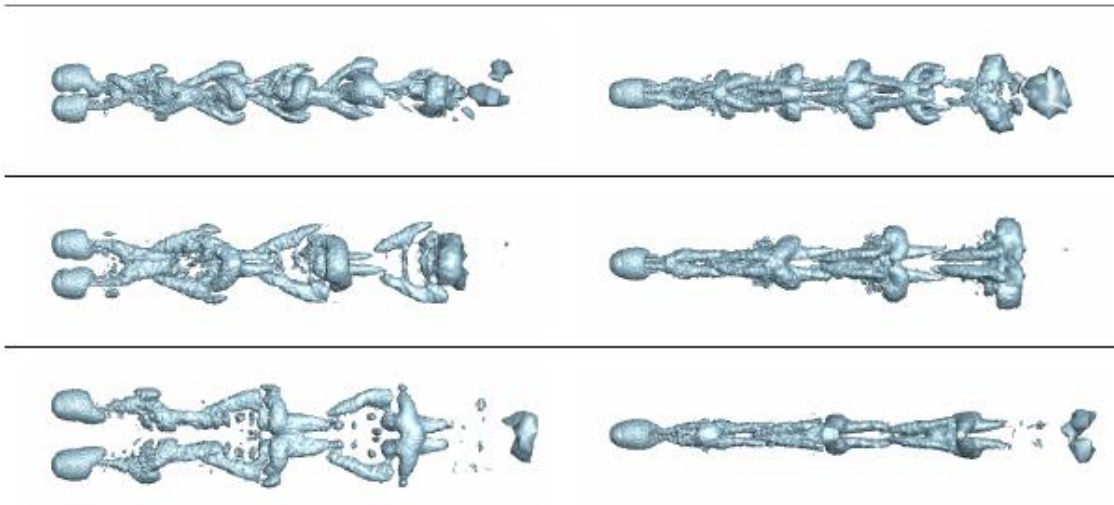
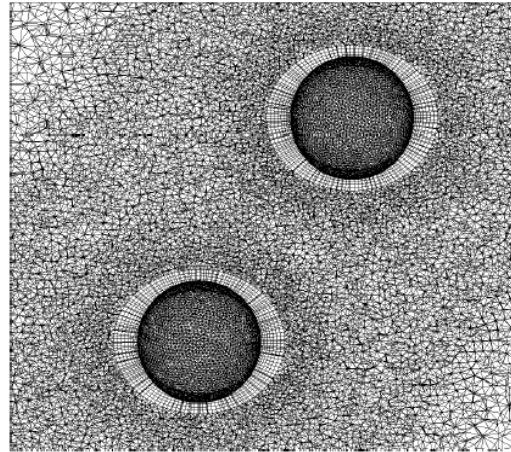
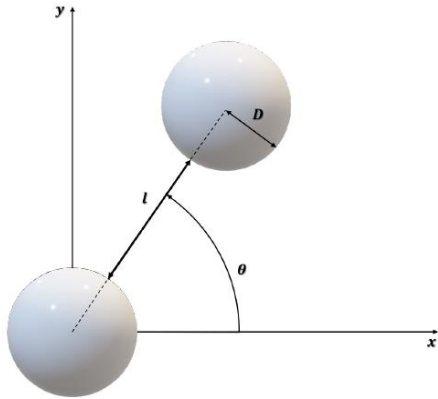
$Re = 300$



$$\Delta C_d = C_d(Re, 1/Fr) - C_d(Re, 0)$$

$$C_d = \frac{F_d}{0.5 \rho_o V_o^2 A}$$

Stratified flow past two spheres

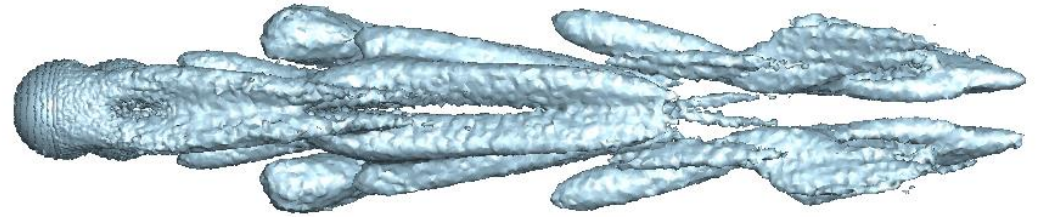
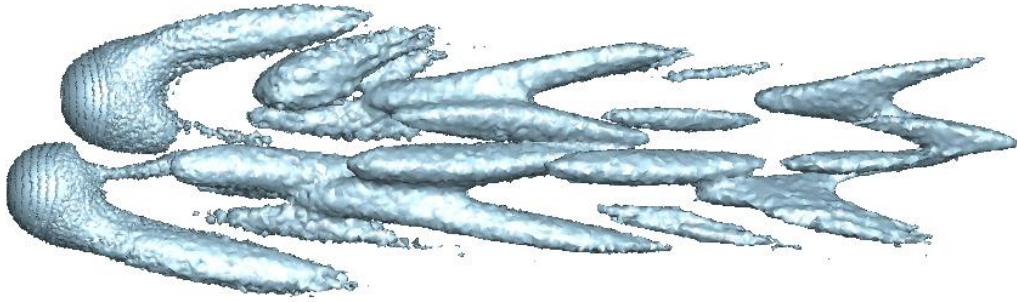


Top view

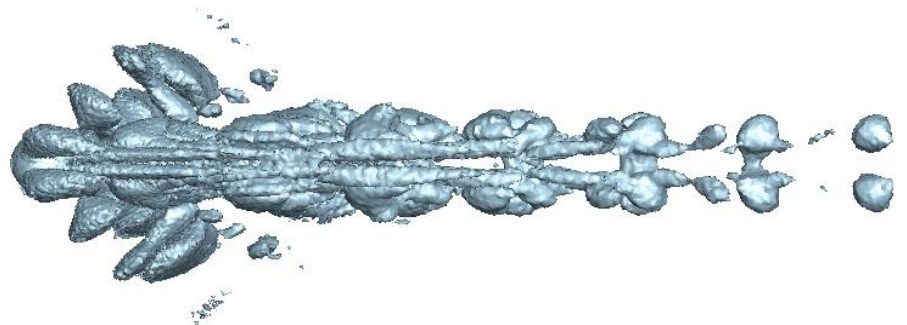
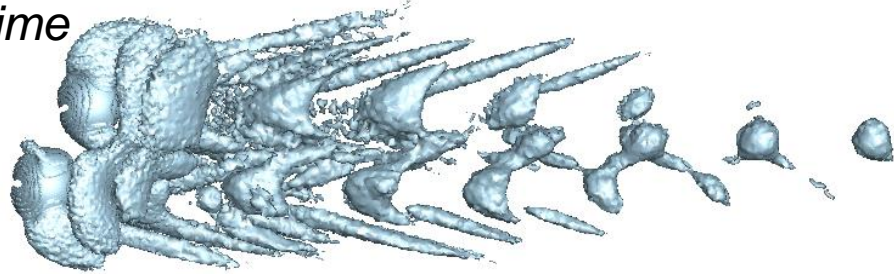
side view

$Re=300$ neutrally stratified

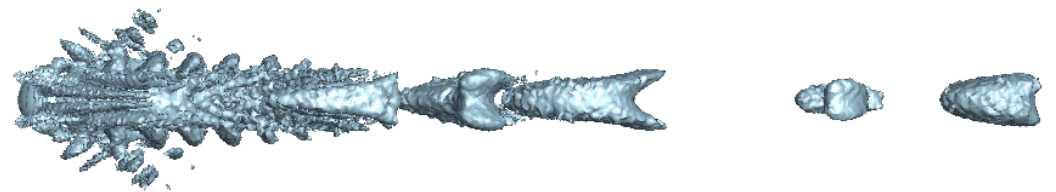
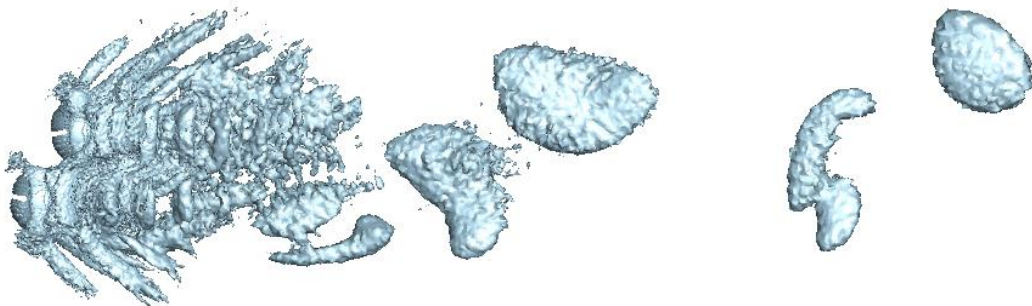
Fr = 1.625 steady-state non-axisymmetric attached vortex regime



Fr=0.625 steady-state lee wave instability regime



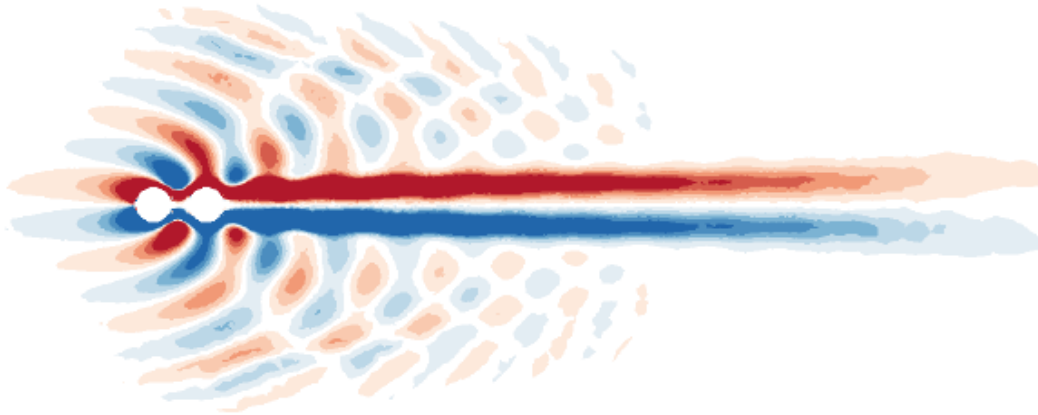
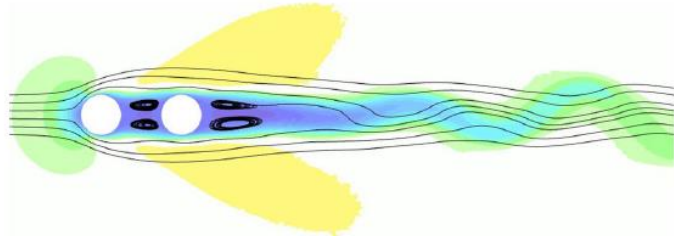
Fr=0.25 two dimensional vortex shedding regime



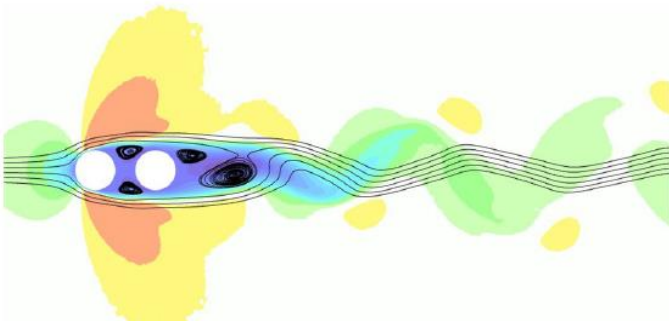
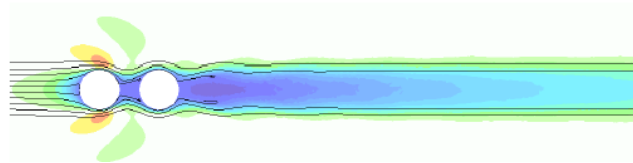
top view

side view

Stratified flow past two spheres



$Fr=0.625$



$Fr=0.25$

top view

side view

Remarks

- The presented NFT methodology can complement established NWP methods.
- Algorithmic parallel scalability is important for future computing architectures. Technical challenges of time to solution & energy efficiency have been partially addressed by a development of specialised preconditioners for the non-symmetric Krylov-subspace solver. For a global atmospheric model, the Jacobi preconditioning on a single mesh provides a substantial speed up.
- General characterisation of flows past spheres confirmed that stratification dominates the flow pattern for $Fr \searrow 0$ and unification of flow patterns. Flow patterns are sensitive to the spheres' relative location. Regimes characterising the flow past a single sphere are also observed for multi-sphere configurations.
- A record of consistent and accurate simulations for NFT MPDATA class of solvers continues to grow. Our recent developments targeted engineering applications and include a development of a LMN solver for strong thermally driven flows and implementation of the dynamic Smagorinski model.