





# Atmospheric gravity waves and their influence on weather and climate: Challenges and new approaches

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# Atmospheric Gravity Waves: Impacts & Issues





Rossby-wave and GW breaking force the Brewer-Dobson circulation





- Rossby-wave and GW breaking force the Brewer-Dobson circulation
- **Downward control:** Middle atmosphere influences tropospheric climate (Haynes et al 1991)







- Rossby-wave and **GW breaking** force the **Brewer-Dobson circulation**
- **Downward control:** Middle atmosphere influences tropospheric climate (Haynes et al 1991)
- BDC under climate change: Impact from GWs (Butchard 2014)







#### Zonal-mean zonal winds (westerlies)



With GW parameterization

Without GW parameterization

Schmidt et al (2006)





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Quasi-Biennial Oscillation (QBO) of zonal-mean zonal wind over the equator:



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Quasi-Biennial Oscillation (QBO) of zonal-mean zonal wind over the equator: Model predictions for QBO under climate change (Schirber et al 2014, Richter et al 2020)



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### Try and resolve everything?



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#### 16. April 2023

## Try and resolve everything? GWD 20°N – 80°N (Polichtchouk et al 2022) No convergence at accessible resolutions 10<sup>1</sup> (e.g. Polichtchouk et al 2022) – -9 km ---4.5 km -1 km 1.4 km explicitly simulated convection Veri 2018110100+36h (h3f7) pressure [hPa] 10<sup>2</sup> -0.40 -0.30 -0.20 -0.10 0.00 GWD [m/s/day]

# **Atmospheric GWs:** Impacts and issues

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## Try and resolve everything?

- No convergence at accessible resolutions (e.g. Polichtchouk et al 2022)
- Model hierarchy needed for conceptual understanding (Held 2005)

1.4 km explicitly simulated convection 2018110100+36h (h3f7)



GWD 20°N – 80°N (Polichtchouk et al 2022)

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# **GW-Mean-Flow Interaction: Theory in a Nutshell**







Short-wavelength GWs in the atmosphere related to geometric optics:

- Atmosphere = medium with refractive properties
- GWs modify the medium
- Two-way interaction described using WKB theory

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#### WKB theory for

- GWs (locally monochromatic or weak amplitude)
- on a slowly varying stratified mean flow u(x, t)

(e.g. Bretherton 1966, Grimshaw 1975, Achatz et al 2017, Achatz 2022):

frequency and wave number connected by dispersion relation

$$\omega = \Omega(\boldsymbol{x}, \boldsymbol{k}, t) = \boldsymbol{k} \cdot \boldsymbol{u}(\boldsymbol{x}, \boldsymbol{t}) \pm \sqrt{\frac{N^2(z)k_h^2 + f^2(\phi)k_z^2}{k_h^2 + k_z^2}}$$

$$\overbrace{\hat{\omega} \text{ (intrinsic frequency)}}^{\text{(intrinsic frequency)}}$$

• GW amplitudes connected by polarization relations, e.g.  $(\hat{b} = g \ \hat{\theta} / \bar{\theta})$ 

$$\widehat{\boldsymbol{u}} = -\frac{i}{mN} \frac{N^2 - \widehat{\omega}^2}{\widehat{\omega}^2 - f^2} \left( \boldsymbol{k}_h \widehat{\omega} - if \boldsymbol{e}_z \times \boldsymbol{k}_h \right) \widehat{\boldsymbol{b}}$$
$$\widehat{\boldsymbol{w}} = \frac{i\widehat{\omega}}{N^2} \widehat{\boldsymbol{b}}$$



**Prognostic equations for weak-amplitude GWs:** 

Spectral wave-action linked to spectral energy density via  $\mathcal{N}(\mathbf{x}, \mathbf{k}, t) = \mathcal{E}(\mathbf{x}, \mathbf{k}, t)/\widehat{\omega}$ 

and satisfies

$$(\partial_t + c_g \cdot \nabla_x + \dot{k} \cdot \nabla_k)\mathcal{N} = D + S$$
  $c_g = \nabla_k \Omega$   $\dot{k} = -\nabla_x \Omega$ 

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D = GW sources and sinks (e.g. wave breaking)

- S = scattering due to GW-GW interactions or interactions with mesoscale balanced motion
- theory from oceanography (Hasselmann 1966, Eden et al 2019) asumes weak mean flows
- neglected in the atmosphere

If D = S = 0 then  $\mathcal{N}$  is conserved along rays (x, k)(t) satisfying  $d_t x = c_g \qquad d_t k = \dot{k}$ 



GW impact on (phase averaged )mean flow, with fluxes from polarization relations:

Horizontal momentum

$$\mathcal{D}_t \boldsymbol{u} = \cdots - \frac{1}{\bar{\rho}} \, \nabla_x \cdot \langle \bar{\rho} \, \boldsymbol{v}' \boldsymbol{u}' \rangle \qquad \langle \bar{\rho} \, \boldsymbol{v}' \boldsymbol{u}' \rangle = \bar{\rho} \int d^3 k \, \widetilde{\boldsymbol{M}}(\boldsymbol{k}) \, \mathcal{N}$$

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Entropy

$$D_t \theta = \dots - \nabla_x \cdot \langle \boldsymbol{u}' \theta' \rangle \qquad \qquad \langle \boldsymbol{u}' \theta' \rangle = \int d^3 k \, \boldsymbol{\Theta}(\boldsymbol{k}) \, \mathcal{N}$$

More conventional (pseudomomentum) approach is

$$D_t \boldsymbol{u} = \cdots - \frac{1}{\bar{\rho}} \, \nabla_x \cdot \int d^3 k \, \boldsymbol{c}_g \boldsymbol{k}_h \, \mathcal{N} \qquad D_t \theta = \cdots - 0$$

but this can lead to errors (Wei et al 2019)

Non-acceleration theorem (Charney & Drazin 1961, Andrews & McIntyre 1978, Achatz 2022): GWs have no impact on a synoptic-scale mean flow if

- they are steady,
- their spatial distribution is horizontally homogeneous, and
- there are no GW sources and sinks (and no wave-wave interactions)





# Numerical Approach: Lagrangian Ray Tracing MS-GWaM

# Numerics: MS-GWaM vs conventional GWP



Non-acceleration theorem (Charney & Drazin 1961, Andrews & McIntyre 1978, Achatz 2022): GWs have no impact on a synoptic-scale mean flow if

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- they are steady,
- their spatial distribution is horizontally homogeneous, and
- there are no GW sources and sinks (and no wave-wave interactions)

#### **Conventional GW parameterizations:**

- Steady state (no transience) Equilibrium profiles assumed, i.e. instantaneous propgation from source to model top
- Single column (1D): no horizontal GW propagation, no horizontal variations of GW energy taken into account
- Rely exclusively on GW breaking

#### **MS-GWaM (Multi-Scale Gravity-Wave Model):**

All three processes taken into account

# Numerics: Conservative Lagrangian Numerics

Without sources and sinks:

$$\left(\partial_t + \boldsymbol{c}_{\boldsymbol{g}} \cdot \nabla_{\boldsymbol{x}} + \dot{\boldsymbol{k}} \cdot \nabla_{\boldsymbol{k}}\right) \mathcal{N} = 0$$

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Phase-space velocity nondivergent  $(\nabla_X \cdot c_g + \nabla_k \cdot \dot{k} = 0)$  $\Rightarrow$  flow is volume preserving

Region of nonzero  ${\mathcal N}$  approximated by rectangular ray volumes

- Ray volumes move with central ray
- Ray volumes change spatial and wavenumber extent in area-preserving manner



20







# **Numerics**: Conservative Lagrangian Numerics





GW packet in Boussinesq flow:

1D, i.e. GW energy and mean flow horizontally homogeneous



time 000 min

# Numerics: MS-GWaM in UA-ICON

MS-GWaM in the

Upper-<u>Atmosphere</u> extension of <u>ICON</u> (DWD/MPI)

 $\Delta x \sim$  160 km (R2B4),  $\Delta z \sim$  1km,  $z_{top} =$  150km



UA-ICON (Borchert et al 2019)

1D framework



Fits well to the current MPI communicator







# Numerics: MS-GWaM in UA-ICON

<u>MS-GWaM</u> in the <u>Upper-Atmosphere</u> extension of <u>ICON</u> (DWD/MPI)

 $\Delta x \sim$  160 km (R2B4),  $\Delta z \sim$  1km,  $z_{top} =$  150km

![](_page_22_Picture_3.jpeg)

UA-ICON (Borchert et al 2019)

1D framework

![](_page_22_Figure_6.jpeg)

Fits well to the current MPI communicator

3D framework

![](_page_22_Figure_9.jpeg)

Requires new MPI communication style for Lagrangian particles

Implementation 3D

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- Cell-based ray-volume handling
- Handover from cell to cell
- MPI parallelization

![](_page_22_Figure_15.jpeg)

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# Numerics: MS-GWaM in UA-ICON

![](_page_23_Picture_1.jpeg)

![](_page_23_Picture_2.jpeg)

#### Sources:

- GWs from convection (based on Choi et al 2011)
  - using latent heat release
  - small-scale waves from convective cells
  - larger-scale waves from mesoscale convective systems
- Background GW source based on Orr et al (2010)
  - Seasonal dependence
  - In each hemisphere horizontally homogeneous
- Orographic GWs parameterized outside MS-GWaM (Lott and Miller 1997)

#### Sink:

Wave breaking due to static instability, using saturation approach (Lindzen 1981, Bölöni et al 2016)

![](_page_24_Picture_0.jpeg)

![](_page_24_Picture_1.jpeg)

10<sup>-5</sup>

0

10<sup>0</sup> 10<sup>0</sup> 90<sup>th</sup> perc 90<sup>th</sup> perc 99<sup>th</sup> perc 99<sup>th</sup> perc Probability of occurrence mean mean occurrence Smooth 2.1 4.3/35% 11.8/ 8% 10<sup>-1</sup> Smooth 1.7 3.8/46% 13.8/13% 10-Mountainous 3.2 5.4/55% 33.7/26% Mountainous 2.3 4.8/58% 28.2/21% 10<sup>-2</sup> 10<sup>-2</sup>  $10^{-3}$  $10^{-3}$ 10-4

60

40

#### Hertzog et al (2012): Vorcore measurements

20

Total momentum flux (mPa)

WRF

40

Total momentum flux (mPa)

![](_page_25_Picture_5.jpeg)

60

![](_page_25_Figure_6.jpeg)

0

20

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![](_page_26_Picture_1.jpeg)

![](_page_26_Picture_2.jpeg)

WRF

## Hertzog et al (2012): Vorcore measurements

10<sup>0</sup> 10<sup>0</sup> 90<sup>th</sup> perc Probability of occurrence 90<sup>th</sup> 99<sup>th</sup> perc 99<sup>th</sup> perc mear perc occurrence mean 2.1 4.3/35% 11.8/8%  $10^{-1}$ Smooth 1.7 3.8/46% 13.8/13% 10-Smooth Mountainous 3.2 33.7/26% Mountainous 2.3 4.8/58% 28.2/21% 10<sup>-2</sup> 10<sup>-2</sup> ≣ 10<sup>0</sup>  $10^{-3}$ TR 20km TR 40km Probability of occurrence TR 60km 10-4 10<sup>-1</sup> TR 80km ST 20km 10<sup>-5</sup> ST 40km ST 60km 20 10<sup>-2</sup> 0 40 ST 80km Total momentum flux (mPa) 10<sup>-3</sup> 10<sup>-4</sup> 10<sup>-5</sup> 10 20 30 40 50 60 0 Bölöni et al (2021): 65°S – 50°S in Oct Absolute zonal momentum flux (mPa)

![](_page_27_Figure_1.jpeg)

![](_page_27_Picture_2.jpeg)

![](_page_27_Figure_3.jpeg)

![](_page_28_Picture_1.jpeg)

![](_page_28_Picture_2.jpeg)

Kim et al (2021): 116°E, 3.5°N in May 1998

![](_page_28_Figure_4.jpeg)

GW source from convection (Son & Chun 2005, Choi & Chun 2011)

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![](_page_29_Picture_2.jpeg)

Gini coefficient: Index for unbalance in distribution

![](_page_29_Figure_4.jpeg)

Cumulative share of people from lowest to highest incomes

![](_page_30_Picture_1.jpeg)

![](_page_30_Picture_2.jpeg)

Gini coefficient: Index for unbalance in distribution

![](_page_30_Figure_4.jpeg)

![](_page_31_Picture_1.jpeg)

![](_page_31_Picture_2.jpeg)

## Kim et al (2021): Gini coefficient for GWMF

![](_page_31_Figure_4.jpeg)

![](_page_32_Picture_0.jpeg)

![](_page_32_Picture_1.jpeg)

# **Effects of Horizontal Propagation**

![](_page_32_Figure_3.jpeg)

![](_page_32_Figure_4.jpeg)

# **Effects of horizontal propagation:** Wave-action budget

![](_page_33_Picture_1.jpeg)

**Spatial wave-action density**  $\mathcal{A} = \int d^3k \mathcal{N}$  satisfies

$$\partial_t \mathcal{A} = -\nabla_h \cdot (\mathbf{c}_{gh} \mathcal{A}) - \partial_z (c_{gz} \mathcal{A}) + D$$

with e.g.  $c_{gh} \mathcal{A} = \int d^3k \ c_{gh} \mathcal{N}$ 

Time mean

$$0 \approx \frac{\Delta \mathcal{A}}{\Delta t} = -\nabla_h \cdot \langle \boldsymbol{c}_{gh} \, \mathcal{A} \rangle - \partial_z \langle \boldsymbol{c}_{gz} \, \mathcal{A} \rangle + \langle D \rangle$$

# **Effects of horizontal propagation:** Wave-action budget

![](_page_34_Picture_1.jpeg)

**Spatial wave-action density**  $\mathcal{A} = \int d^3k \mathcal{N}$  satisfies

$$\partial_t \mathcal{A} = -\nabla_h \cdot (\mathbf{c}_{gh} \mathcal{A}) - \partial_z (c_{gz} \mathcal{A}) + D$$

with e.g. 
$$c_{gh} \mathcal{A} = \int d^3k \ c_{gh} \mathcal{N}$$

Time mean

![](_page_34_Figure_6.jpeg)

60°N 30°N

0°

30°5 60°S

June  $z \approx 40 \text{km}$ (Völker et al 2023, in prep.)

![](_page_34_Figure_8.jpeg)

# **Effects of horizontal propagation:** Horizontal distribution GW mom.flux

![](_page_35_Picture_1.jpeg)

![](_page_35_Picture_2.jpeg)

GW momentum flux November (snapshot) at two altitudes (Völker et al 2023, in prep.)

![](_page_35_Figure_4.jpeg)

![](_page_35_Figure_5.jpeg)

# **Effects of horizontal propagation:** GW mom.flux & mean winds

![](_page_36_Picture_1.jpeg)

June 1994 Southern Hemisphere **GW meridional momentum flux & mean zonal wind** (Völker et al 2023, in prep.)

![](_page_36_Figure_3.jpeg)

# **Effects of horizontal propagation:** GWMF intermittency

## Kim et al (2023, in prep.)

![](_page_37_Figure_2.jpeg)

![](_page_37_Figure_3.jpeg)

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# **Effects of horizontal propagation:** Zonal-mean zonal wind

![](_page_38_Picture_1.jpeg)

![](_page_38_Picture_2.jpeg)

#### Völker et al (2023, in prep.)

![](_page_38_Figure_4.jpeg)

**Effects of horizontal propagation:** QBO

![](_page_39_Picture_1.jpeg)

![](_page_39_Picture_2.jpeg)

**Quasi-Biennial Oscillation:** zonal-mean zonal wind  $5^{\circ}S - 5^{\circ}N$  (Kim et al 2023, in prep.)

![](_page_39_Figure_4.jpeg)

Year

Quasi-Biennial Oscillation: zonal-mean zonal wind 5°S – 5°N (Kim et al 2023, in prep.)

![](_page_40_Figure_1.jpeg)

**Effects of horizontal propagation:** QBO

![](_page_40_Picture_4.jpeg)

transient 3D

![](_page_40_Picture_5.jpeg)

**Effects of horizontal propagation:** QBO

![](_page_41_Picture_1.jpeg)

![](_page_41_Picture_2.jpeg)

**Quasi-Biennial Oscillation:** zonal-mean zonal wind  $5^{\circ}S - 5^{\circ}N$  (Kim et al 2023, in prep.)

![](_page_41_Figure_4.jpeg)

**Effects of horizontal propagation:** QBO

![](_page_42_Picture_1.jpeg)

![](_page_42_Picture_2.jpeg)

Quasi-Biennial Oscillation: zonal-mean zonal wind  $5^{\circ}S - 5^{\circ}N$  (Kim et al 2023, in prep.)

![](_page_42_Figure_4.jpeg)

![](_page_43_Picture_0.jpeg)

![](_page_43_Picture_1.jpeg)

# Is it affordable?

![](_page_44_Picture_1.jpeg)

![](_page_44_Picture_2.jpeg)

	UA-ICON / MS-GWaM (3D)
$\Delta x/\text{km}$	160
$\Delta z/km$	0.02 (BL) – 0.7 (ST) – 4.5 (top)
relative increase comp. time (compared to ICON with classic GWP - Orr et al 2010)	10-40

UA-ICON / MS-GWaM is

• more expensive than UA-ICON with classic GW parameterization, but

![](_page_45_Picture_1.jpeg)

![](_page_45_Picture_2.jpeg)

	UA-ICON / MS-GWaM (3D)	UA-ICON GW resolving
$\Delta x/\mathrm{km}$	160	5
$\Delta z/\mathrm{km}$	0.02 (BL) – 0.7 (ST) – 4.5 (top)	0.02 (BL) – 0.2 (ST) – 0.2 (top)
relative increase comp. time (compared to ICON with classic GWP - Orr et al 2010)	10-40	240000

#### UA-ICON / MS-GWaM is

- more expensive than UA-ICON with classic GW parameterization, but
- considerably cheaper than a wave-resolving set-up of UA-ICON

# **Summary & Discussion**

![](_page_46_Picture_1.jpeg)

![](_page_46_Picture_2.jpeg)

- Subgrid-scale processes still present major challenges to the reliability of climate simulations
- Gravity waves (GW) are a corresponding phenomenon
- Both transient dynamics and horizontal propagation not represented in present-day parameterizations
- MS-GWaM in UA-ICON is 1st prognostic GW model to simulate these effects
- Corresponding differences are leading order, e.g.
  - Intermittency
  - Horizontal distributions
  - QBO
- MS-GWaM is more expensive than classic GW parameterizations but much cheaper than wave-resolving simulations
- Even when we will be able to resolve all processes even in climate simulations, conceptional models with a solid theoretical basis will remain essential for our gain of understanding

## Literature

![](_page_47_Picture_1.jpeg)

![](_page_47_Picture_2.jpeg)

#### Book:

Achatz, U., 2022: Atmospheric Dynamics. Springer, Berlin

![](_page_47_Picture_5.jpeg)

#### **Journal Papers:**

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- Kim, Y.-H., G. Bölöni, S. Borchert, H.-Y. Chun, and U. Achatz, 2021: Toward transient subgrid-scale gravity wave representation in atmospheric models. Part II: Wave intermittency simulated with convective sources. J. Atmos. Sci., **78**, 1339–1357
- Muraschko J, Fruman M, Achatz U, Hickel S, Toledo Y. 2015: On the application of WKB theory for the simulation of the weakly nonlinear dynamics of gravity waves. *Quart. J. R. Met. Soc.* **141**, 676–697.
- Wei, J., Bölöni, G., and U. Achatz 2019: Efficient modeling of the interaction of mesoscale gravity waves with unbalanced large-scale flows: Pseudomomentum-flux convergence versus direct approach, *J. Atmos. Sci.*, **76**, 2715–2738
- Wilhelm, J., Akylas, T.A., Bölöni, G., Wei, J., Ribstein, B., Klein, R., and U. Achatz, 2018: The interaction between meso- and sub-mesoscale gravity waves. *J. Atmos. Sci.*, **75**, 2257-2280

MS-GWaves: https://ms-gwaves.iau.uni-frankfurt.de/